1 2 3 Ozonesonde Quality Assurance: The JOSIE-SHADOZ (2017) Experience 4 5 6 Anne M. Thompson<sup>1\*</sup>, Herman G. J. Smit<sup>2</sup>, Jacquelyn C Witte<sup>1,3</sup>, Ryan M. Stauffer<sup>1,4</sup>, Bryan J. 7 8 Johnson<sup>5</sup>, Gary Morris<sup>6</sup>, Peter von der Gathen<sup>7</sup>, Roeland Van Malderen<sup>8</sup>, Jonathan Davies<sup>9</sup>, Ankie Piters<sup>10</sup>, Marc Allaart<sup>10</sup>, Françoise Posny<sup>11</sup>, Rigel Kivi<sup>12</sup>, Patrick Cullis<sup>5,13</sup>, Nguyen Thi 9 Hoang Anh<sup>14</sup>, Ernesto Corrales<sup>15</sup>, Tshidi Machinini<sup>16</sup>, Francisco R. da Silva<sup>17</sup>, George Paiman<sup>18</sup>, 10 Kennedy Thiong'o<sup>19</sup>, Zamuna Zainal<sup>20</sup>, George B. Brothers<sup>21,22</sup>, Katherine R. Wolff<sup>3,22</sup>, Tatsumi 11 Nakano<sup>23</sup>, Rene Stübi<sup>24</sup>, Gonzague Romanens<sup>24</sup>, Gert J. R. Coetzee<sup>16</sup>, Jorge A. Diaz<sup>15</sup>, Sukarni 12 Mitro<sup>18</sup>, Maznorizan 'bt Mohamad<sup>20</sup>, Shin-Ya Ogino<sup>25</sup> 13 14 15 12 August 2018 16 <sup>1</sup> NASA Goddard Space Flight Center, Greenbelt, MD 17 18 <sup>2</sup> Institute of Chemistry and Dynamics of the Geosphere: Troposphere, Research Centre Jülich, 19 Jülich, Germany <sup>3</sup> Science Systems and Applications Inc., Lanham, MD 20 <sup>4</sup> Universities Space Research Association, Columbia, MD 21 22 <sup>5</sup> NOAA Earth System Research Laboratory, Global Monitoring Division, Boulder, CO 23 <sup>6</sup> St. Edwards University, Natural Sciences, Austin, TX 24 <sup>7</sup> Alfred Wegener Institute, Potsdam, Germany 25 <sup>8</sup> Royal Meteorological Institute of Belgium, Brussels, Belgium <sup>9</sup> Environment and Climate Change Canada, Toronto, Ontario, Canada 26 <sup>10</sup> Royal Dutch Meteorological Institute, de Bilt, The Netherlands 27

- <sup>11</sup> Laboratoire de l'Atmosphère et des Cyclones, UMR8105 (Université, Météo-France, CNRS), La 28
- 29 Réunion, France
- <sup>12</sup> Finnish Meteorological Institute, Sodankylä, Finland 30
- <sup>13</sup> Cooperative Institute for Research in Environmental Sciences, Boulder, CO 31
- <sup>14</sup> Vietnam Meteorological Hydrological Administration, Ha Noi, Vietnam 32
- <sup>15</sup> University of Costa Rica, San José, San Pedro, Costa Rica 33
- <sup>16</sup> South African Weather Service, Pretoria, South Africa 34
- <sup>17</sup> Laboratory of Environmental and Tropical Variables, Brazilian Institute of Space Research, Natal, 35
- 36 **Brazil**
- 37 <sup>18</sup> Meteorological Service of Suriname, Paramaribo, Surinam
- <sup>19</sup> Kenyan Meteorological Department, Nairobi, Kenya 38
- <sup>20</sup> Malaysian Meteorological Department, Atmospheric Science and Cloud Seeding Division. 39
- 40 Petaling Jaya, Selangor, Malaysia
- 41 <sup>21</sup> CHEMAL, Wallops Is., VA
- <sup>22</sup> NASA Wallops Flight Facility, Wallops Is., VA 42
- <sup>23</sup> Japan Meteorological Agency, Tokyo, Japan 43
- <sup>24</sup> MeteoSwiss, Paverne, Switzerland 44
- <sup>25</sup> Japan Agency for Marine-Earth Science and Technology, Department of Coupled Ocean-45
- 46 Atmosphere-Land Processes Research, Yokosuka, Japan

\*Corresponding Author: anne.m.thompson@nasa.gov; 301-614-5905; NASA/GSFC 48

49 50

54

55

56

57

58

59

60

61

62

63

64

65

66

67

68

69

70

71

72

73

74

Abstract. The ozonesonde is a small balloon-borne instrument that is attached to a standard radiosonde to measure profiles of ozone from the surface to 35 km with ~100-m vertical resolution. Ozonesonde data constitute a mainstay of satellite calibration and are used for climatologies and analysis of trends, especially in the lower stratosphere where satellites are most uncertain. The electrochemical-concentration cell (ECC) ozonesonde has been deployed at ~100 stations worldwide since the 1960s, with changes over time in manufacture and procedures, including details of the cell chemical solution and data processing. As a consequence, there are biases among different stations and discontinuities in profile time-series from individual site records. For 22 years the Jülich [Germany] Ozone Sonde Intercomparison Experiment (JOSIE) has periodically tested ozonesondes in a simulation chamber designated the World Calibration Centre for Ozonesondes (WCCOS) by WMO. In October-November 2017 a JOSIE campaign evaluated the sondes and procedures used in SHADOZ (Southern Hemisphere Additional Ozonesondes), a 14-station sonde network operating in the tropics and subtropics. A distinctive feature of the 2017 JOSIE was that the tests were conducted by operators from eight SHADOZ stations. Experimental protocols for the SHADOZ sonde configurations, which represent most of those in use today, are described, along with preliminary results. SHADOZ stations that follow WMO-recommended protocols record total ozone within 3% of the JOSIE reference instrument. These results and prior JOSIEs demonstrate that regular testing is essential to maintain best practices in ozonesonde operations and to ensure high-quality data for the satellite and ozone assessment communities. **Capsule:** Data from ozonesondes form a backbone of satellite algorithms and monitoring stratospheric ozone recovery. The ozonesonde community regularly evaluates sonde procedures and instrumentation, as in this experiment featuring operators from the tropical SHADOZ

#### network.

76

77

78

79

80

81

82

83

84

85

86

87

88

89

90

91

92

93

94

95

96

97

98

75

### **JOSIE History and Background**

The periodic ozone assessments sponsored by WMO/UNEP (1991; 1995; 2011; 2015) and related studies have long recognized the role of ozonesondes in the suite of global observations because sondes are the only technique practical for in-situ monitoring of profiles. The sonde instrument is easy to deploy in remote locations and is relatively inexpensive. Sondes operate in both troposphere and stratosphere (Sidebar 1) and in clouds, precipitation and periods of darkness. Most important, as they ascend, ozonesondes measure ozone with an effective resolution of 100-150 m, far better than satellites. Indeed, sondes, like the ground-based networks of lidar, Dobson and other spectrometers, constitute an essential component of satellite calibration and cross-calibration (Fishman et al., 2008; Hubert et al., 2016; Steinbrecht et al., 2017; Tarasick et al., 2018). The vertical structure of ozone as measured at a typical tropical station appears in Sidebar 1, along with background on ozone in the atmosphere. Although dozens of stations began launching ozonesondes in the 1970s and 1980s, the concepts of standardizing and testing instruments in a coordinated network, did not evolve until the 1990s (Mohnen, 1996; Melamed et al., 2015). This was the period when both the Jülich Ozone Sonde Intercomparison Experiment (JOSIE) and Southern Hemisphere Additional Ozonesondes (SHADOZ) project began.

#### [Insert Sidebar 1 Here]

Over 50 years of ozonesonde data-taking, there have been several instrument designs. Furthermore, as instruments have changed and preparation and data-processing techniques have evolved over time, time series of data from individual stations often display discontinuities and gaps that lead to inhomogeneous data records. Thus, the reliability of ozonesonde trends was

questioned in some of the earlier ozone assessments (WMO/UNEP 1991; 1995; SPARC/IOC/GAW, 1998) (See Acronym List).

99

100

101

102

103

104

105

106

107

108

109

110

111

112

113

114

115

116

117

118

119

120

121

122

Two approaches have been used to address these deficiencies. First, evaluations of ozonesonde types in a controlled laboratory environment were undertaken in the 1990s, a process that continues periodically to this day. Second, in a similar manner, by testing different sonde preparation methods and protocols for data recording and processing, a set of standard operating procedures (SOP; Smit et al., 2014) was developed through consensus with the ozonesonde research community. Finally, there are recommended methods for reprocessing long-term records compromised by inhomogeneities (Smit et al., 2012, Deshler et al., 2017). The need to have recommended instruments and procedures for emerging WMO/GAW stations in the 1990s provided a framework for the first intercalibration and intercomparisons of existing ozonesonde types. In order to assess the performance of the various ozonesonde instrument types used within GAW, the environmental simulation chamber (ESC) at the Forschungszentrum Jülich (FZJ, Germany) was established as the World Calibration Centre for Ozone Sondes (WCCOS) in 1996. The chamber enables control of pressure, temperature, and ozone concentration as it simulates flight conditions of ozone soundings up to an altitude of 35 km (Smit et al., 2000). This controlled environment and comparison of the ozonesonde profiles with an accurate UV-photometer as a reference (Proffitt and McLaughlin, 1983) are essential requirements for addressing instrument issues that arise from field and laboratory operations. The initial JOSIE, performed in 1996 (Smit and Kley, 1998), was the first GAW activity directed toward implementing a global quality assurance plan for ozonesondes in routine use. By now, JOSIE experiments have provided over twenty years of ozonesonde data quality assurance to the larger atmospheric research and remote sensing communities. JOSIE-1996 was

attended by eight laboratories from seven countries representing the major types of ozonesondes:

Electrochemical Concentration Cell (ECC) sondes of two manufacturers, the Brewer/Mast sonde (BM-original), the Indian sonde (a modified BM-type), and the Japanese Meisei sonde (KC79). JOSIE-1996 revealed important information not only about ozonesonde performance but also the influence of operating procedures for sonde preparation and data correction that often varied among the participating laboratories. The succession of JOSIE campaigns (**Table 1**) has shown that there is an on-going need to evaluate ozonesondes because the instruments, preparation procedures, and/or the sensing solutions are modified, often inadvertently, over time. Routine testing of newly manufactured ozonesondes on a regular basis coupled with better standardization of operating procedures help ensure more confidence in the data itself as well as trends calculated from the data.

The overall objective of WCCOS and the JOSIE series of experiments has been the establishment of a facility for ozonesonde quality assurance (QA) that can be used by sonde manufacturers and the research community. Instrumental performance of sondes from different manufacturers is tested through comparison of profiling capabilities with a standard ozone profile that simulates a typical ascent in polar, mid-latitude or tropical conditions. Regular evaluation of procedures and methods at long-term ozone sounding stations with a single ozone reference instrument ensures the traceability and consistency of the records.

Over time, the SOP have been established and updated as needed. The first major SOP documentation appeared as a WMO/GAW Report (#201; See Smit and ASOPOS, 2014) with major contributions from prior reports and Smit et al. (2007). GAW 201 was also based on field tests of the major sonde types used in the JOSIEs up through 2009. A gondola of 18 instruments was flown along with same UV-photometer used in JOSIE-2000 as reported in Deshler et al. (2008).

#### **SHADOZ and Unresolved Sonde Issues**

147

148

149

150

151

152

153

154

155

156

157

158

159

160

161

162

163

164

165

166

167

168

169

170

The SHADOZ network began in 1998 as an international partnership to enhance the number of tropical ozone soundings from operational stations (Thompson et al., 2003a,b; 2004; 2007; 2011). SHADOZ uses ECC ozonesondes that, over time, have been coupled with a variety of radiosondes (Table 2). A history of ozonesonde-radiosonde pairings used at SHADOZ sites appear in archival papers (Thompson et al., 2003a,b; Thompson et al., 2007; Witte et al., 2017). At the time SHADOZ began, all known operational stations were in the southern hemisphere, but gradually northern hemisphere stations joined: Kuala Lumpur, Paramaribo, Costa Rica; Hanoi, and Hilo. The 14 long-term stations, defined as operating at least a decade during SHADOZ, appear in **Fig. 1.** More than 7000 sets of ozone and pressure-temperature-humidity profiles from SHADOZ are available at the website: https://tropo.gsfc.nasa.gov/shadoz. Periodic evaluations of SHADOZ data have examined three parameters. First, total column ozone (TCO) from the sonde, with an appropriate extrapolation above balloon burst, e.g., McPeters and Labow, (2012), is compared to TCO from co-located ground-based instruments (Brewer, Dobson, SAOZ) and satellite overpasses. Second, stratospheric profiles are compared to satellite overpass ozone profiles from instruments like SAGE II (to 2005), SBUV (entire record, 1998-2016) or Aura's MLS (2005-). Third, for the tropical stations (generally within 18° latitude of the equator), stratospheric column ozone and profiles are compared. The tropical TCO is typically constant to within 3-5 DU (Dobson Units), so measurement biases from station to station can be identified (Thompson et al., 2017). The first three years of SHADOZ TCO compared to the EP/TOMS satellite TCO disagreed by ~8% on average, with a number of stations displaying a discrepancy of greater than 10%; the sonde TCO was usually lower than the satellite (or ground-based instrument). After the JOSIE-2000 campaign (Smit et al., 2007), in which the instruments and techniques used at all the

SHADOZ stations were tested, several stations changed their sensing solution type (SST), resulting in reduced offsets (Thompson et al., 2007). Further changes in sonde preparation procedures and subsequent reprocessing of the data, both in accordance with WMO/SPARC/ IOC/NDACC guidelines (Smit and O3S-DQA, 2012; Smit and ASOPOS, 2014), brought TCO for 12 of 14 stations to within 2% of TCO from three BUV-type satellites (EP/TOMS, OMI and OMPS) operating over the 1998-2016 period (Thompson et al., 2017); the remaining two stations show TCO data averaging within 5% of the satellite TCO. These improvements derive from the application of "transfer functions" that relate a profile from each instrument—SST combination to data from the standard reference. Each profile in a time-series is examined for possible correction (Witte et al., 2017; 2018).

Although the reprocessing of prior SHADOZ data has greatly reduced systematic variations in the record, JOSIE-SHADOZ was designed to address several outstanding issues. First, transfer functions determined by Deshler et al. (2017) are used to homogenize SHADOZ readings that are taken with different SST and/or instruments. This includes the 1%, KI, 0.1% buffer SST used at stations supported by NOAA since the mid-2000s (Sterling et al., 2018). Second, a few stations in SHADOZ changed SST unintentionally and introduced discontinuities in station time-series (Thompson et al., 2017; Witte et al., 2017; 2018). Finally, several stations employing a given sonde type show sharp discontinuities after 2014 that appear to originate with changes in manufacture (Sterling et al., 2018; Thompson et al., 2017).

#### [Insert Sidebar 2 here]

#### **JOSIE-SHADOZ-2017 Goals**

- Similar to prior JOSIE campaigns, the major objectives of JOSIE-SHADOZ are:
- 194 1. Evaluate ozonesonde instrument performance, specifically the pump and sensor as delivered

- by the ECC-sonde manufacturer. Most of the SHADOZ stations operate with WMO-recommended solutions and preparation and calibration procedures that allow the experimenters to update typical performance of the instruments relative to the Ozone Photometer (OPM) reference instrument (Proffitt and McLaughlin, 1983).
- 2. Evaluate current preparation and operating procedures of each SHADOZ station. Unlike prior JOSIE experiments, in 2017 personnel representing the practices of all currently operating SHADOZ stations participated (Tables 2 and 3; see Sidebar 2). In most cases the operators supplied solutions as prepared at their home institution. In the first part of the JOSIE-2017, the operators followed their standard practice for pre-conditioning sondes and for "day of flight" prior to simulation in the ESC. The goal was to understand the existing ozone profiles archived in SHADOZ by reproducing current practices, techniques, and solutions at each participating station as closely as possible.
  - 3. Evaluate the current WMO recommended SOP. Specific instrumental aspects examined in these tests were details of pre-conditioning, background current, response time, pump flow efficiency, and SST. In addition to two WMO-recommended SST, two alternatives, one of which is employed at several SHADOZ stations, were included in the tests.

#### **The Ozonesonde Design**

The electrochemical concentration cell (ECC) ozonesonde uses a chemical reaction measured inside a pair of cells that is displayed schematically in **Fig. 3a.** As the sonde rises in the atmosphere (and during the laboratory calibration phase), air is pulled through the intake tube (left in **Fig. 3a**) and pushed into the cathode cell by means of a small pump. The pump maintains positive pressure as the air is sampled; the flow rate is measured during pre-flight calibration. The second cell (anode) is filled with a saturated version of the cathode solution and

219 is located adjacent to the cathode, with an ion bridge separating the two cells. The reacting

chemical, oxidized by the ozone molecule, is dissolved potassium iodide (KI). The sensing

solution is maintained at a neutral pH with the addition of the paired phosphates (NaH<sub>2</sub>PO<sub>4</sub>•H<sub>2</sub>O/

- Na<sub>2</sub>HPO<sub>4</sub>•12H<sub>2</sub>O). The ozone partial pressure is calculated by the following equation (taken
- 223 from Witte et al. 2018),

224 
$$R_{O3} = 4.307 \cdot 10^{-2} \frac{\left(I_{M} - I_{B}\right)T_{R}}{Y_{R}F_{R}h_{C}}$$

- 225 where
- $P_{O3} = Ozone partial pressure, mPa$
- 227  $I_M = Cell current, \mu A$
- 228  $I_B$  = Cell background current,  $\mu A$
- $T_P = Ozonesonde pump temperature, K$
- 230  $\Phi_P = \text{Pump flow-rate, ml/s}$
- 231  $\Psi_P$  = Pump flow efficiency, unitless
- $\eta_C$  = Conversion efficiency which is generally assumed to be 1.

233

235

236

237

The pump flow efficiencies,  $\Psi_P$ , take into account the buffering of the solution, depending on the

solution recipe, and mechanical degradation of the pump at low pressures (< 100 hPa). The

volume mixing ratio is computed from the ratio of the ozone partial pressure (P<sub>O3</sub>) to the ambient

pressure determined from the radiosonde attached to the ozonesonde container as the two

- instruments ascend into the stratosphere (**Fig. 3b**). The typical ascent rate is 5 m/s.
- From the large body of SHADOZ data as well as instruments in the field and prior lab

intercomparisons it is known that the two major sources of systematic error are the manufacture

of the instrument and the composition of the KI and/or buffers in the SST (Smit et al., 2007).

243

244

245

246

247

248

249

250

251

252

253

254

255

256

257

258

259

260

261

262

263

264

Random sources of error include operator handling and changing conditions in the station calibration unit. Calibration practices and the method of data-processing can also lead to systematic differences among station profiles (Johnson et al., 2002; Deshler et al., 2008; 2017). In JOSIE-SHADOZ two types of protocols investigated these issues. The first five of ten tests in each session were carried out with the operators using their own solutions and preparation technique. We refer to this as SHADOZ SOP (Standard Operating Procedure). In the second set of tests, uniform calibration and preparation procedures were followed using JOSIE-prepared solutions, hereafter referred to as the JOSIE SOP. Unified data collection by the Data Acquisition System (DAS) eliminates variations due to operator data-processing. General Operations during JOSIE-SHADOZ (2017). The JOSIE-SHADOZ 2017 campaign took place at the World Calibration Centre for Ozone Sondes (WCCOS) at the Research Center Jülich (FZJ) in the Institute of Energy and Climate Research: Troposphere (IEK-8), Jülich, Germany. Ozonesonde pre-conditioning test units and the ECC instruments were provided by FZJ from a pool of loaned supplies. Participants were split into two groups (Table 3), each of four teams operating ozonesondes of the type used in SHADOZ (Table 2). Each group participated in a 12-day intercomparison campaign. Session No. 1 took place from 9 to 20 October 2017; Session No. 2 took place from 23 October through 3 November 2017. Each session consisted of ten simulation experiments with all four participant sondes being "flown" simultaneously in the chamber (see **Sidebar 3**) to an effective altitude of ~35 km. The overall protocol for each campaign was similar but the second session tested two "JOSIE SSTs" (Table 4). During the SHADOZ SOP (first five simulations) participants used their own zero-air filter, solutions, and preparation procedures. During the JOSIE SOPs the lab provided a single source

265 of high-quality zero-air, a common SST, and common operating procedures that all teams 266 followed. Data were collected by the DAS of the WCCOS test chamber. 267 [Insert Sidebar 3 here] 268 Because JOSIE-SHADOZ 2017 was focused on questions about SHADOZ operations, all 269 the chamber runs simulated tropical sounding conditions (Fig. 4). The test profiles described in 270 Fig. 4 and Table 4 represent three typical tropical profiles, one that is unpolluted throughout the 271 troposphere with very low ozone near the tropopause and two with higher levels of ozone in the 272 free troposphere and near the tropopause. 273 Four SST recipes were tested. All sonde data were processed by using a constant 274 background current correction. Total ozone column normalization was not applied. The 275 solutions, with references, follow: 276 1. SHADOZ 1.0. The WMO-recommended SOP (Smit et al., 2012) for use with the 277 Science Pump (SPC) instrument and is referred to as SST 1.0%-Full Buffer: 278 1% KI + Full-Buffer & KBr as described by Komhyr (1986) Cathode: 279 Anode: Cathode solution with saturated KI 280 Pump flow efficiency factors (PEF): Komhyr (1986) 281 2. SHADOZ 0.5. The WMO-recommended SOP (Smit et al., 2012) for use with the ENSCI 282 instrument is referred to as SST 0.5%-Half Buffer: 283 Cathode: 0.5% KI + Half of the Buffer & KBr as described by Komhyr et al. (1995) 284 Anode: Cathode solution with saturated KI 285 PEF: Komhyr et al. (1995) 286 3. JOSIE 1.0.1. Solution developed by NOAA for use with ENSCI sondes that has been 287 employed at Fiji, Samoa, Costa Rica, and Hilo stations since the late 2000's. The

formulation is SST 1.0%-1/10<sup>th</sup> Buffer:

1% KI+ 1/10<sup>th</sup> Buffer, KBr as described by Komhyr (1986) 289 Cathode: 290 Anode: Cathode solution with saturated KI 291 PEF: New constants derived from recent pumpflow measurements made by 292 Nakano (2017, private communication). 4. JOSIE 2.0.1. This variation on JOSIE 1.0.1 was used to test if ozone response in the 293 294 tropopause and stratosphere regions is improved by doubling the KI concentration: 2% KI + 1/10<sup>th</sup> Buffer, KBr as described by Komhyr (1986) Cathode: 295 296 Cathode solution with saturated KI Anode: 297 PEF: New constants derived from recent pumpflow measurements made by 298 Nakano (2017, private communication).

299

300

301

302

303

304

305

306

307

308

309

310

311

#### **Preliminary Results**

Preliminary data are used to answer three questions. (1) What is the accuracy of ozone readings throughout the profile for each sonde-SST combination tested in the ESC? This is answered by comparing both the ozone partial pressure profiles measured by the sonde with the OPM and column-integrated ozone from the sondes with the OPM. For the latter, TCO and segments for troposphere, stratosphere and the tropopause transition layer (TTL) in between the stratosphere and troposphere are computed. (2) How do profiles and column segments from sondes prepared with the SHADOZ SOP compare to those prepared with the JOSIE SOP? (3) What differences are observed when the same instrument type is prepared with different SST or when different instruments use the same SST? Differences are expected based on prior JOSIE results and field tests.

SHADOZ SOP. Fig. 5 displays raw data from eight SHADOZ participants. The OPM

312	measurements are represented by the black dashed lines: Fig. 5a shows the data for a simulation
313	in Session 1 (No. 171) and <b>Fig. 5b</b> for a simulation in Session 2 (No. 182). The fundamental unit
314	in the tests is lapsed time; quoted "altitudes" are approximate. There is some arbitrariness in
315	designating the TTL, with lower-mid-troposphere below and mid to upper stratosphere above.
316	We adopt a TTL at 2200-3800 s (~12-18 km) when analyzing the test results. In this region the
317	signal-to-noise ratio is low, and therefore the uncertainty, is highest (Witte et al., 2017).
318	In Fig. 5a the ozone partial pressures are very small throughout the "troposphere" and up
319	to ~3500 s or ~17.5 km. This profile simulates a near-zero-ozone tropopause, mimicking
320	western Pacific profiles (Kley et al., 1996; Thompson et al., 2012; Rex et al., 2014; Newton et
321	al., 2016), where SNR in ozone readings is often low. In Fig. 5b ozone partial pressure
322	throughout the tropospheric profile is higher, representing stations influenced by biomass
323	burning pollution in the lower-mid troposphere (Thompson et al., 1996; Jensen et al., 2012). The
324	ozone transition near the tropopause and in the lower stratosphere in Simulation No. 182 (Fig.
325	5b) lacks the sharp gradient intentionally generated in Fig. 5a. The pattern in Fig. 5b resembles
326	that of SHADOZ stations that exhibit gradual ozone transitions in the TTL, e.g., Ascension,
327	Natal and Nairobi. Their upper tropospheric and TTL cross-sections and their contributions to
328	the zonal wave-one in tropical ozone are summarized in Thompson et al. (2003b; 2011; 2017).
329	The OPM TCO in Fig. 5a is 282 DU. The TCO from the four participants in Session 1
330	are all higher than the OPM by 3-26 DU (up to 9%). The OPM TCO in <b>Fig. 5b</b> is 334 DU. The
331	TCO from the four participants in Session 2 are all equal to or higher than the OPM, with the
332	largest offset 23 DU (7%) higher. Columns 2 and 3 in <b>Table 5</b> list the corresponding TCO
333	fractions for all 8 participants relative to the OPM.
334	The means of five simulations for all eight participants, expressed as absolute and
335	percentage differences from the OPM and based on their SHADOZ SOP are displayed in Fig. 6.

336 The shapes of the mean profiles are broadly similar with the sonde partial pressures (relative to 337 the OPM, Fig. 6a) overlapping throughout the troposphere and TTL (to 3500s). In the 338 stratosphere (above 4000 s, ~20 km) differences are much larger. The fractional differences are 339 smaller in the stratosphere (Fig. 6b), however, because the ozone partial pressure peaks at over 340 20 mPa (Fig. 5). The relative differences with the OPM are largely within + 10% of the OPM 341 (zero-line in **Fig. 6b**) throughout the lower to mid-troposphere (0-2000 s, up to 10-12 km). 342 Around 2000 s, there is an inflection, with the offsets all turning more negative. The largest 343 relative differences occur within the upper troposphere (UT) and TTL (equivalent to 2500-3500 344 s, 13-18 km), exceeding 5% on average for all the stations. For participant nos. 4 and 5 the mean 345 relative differences exceed -20%. Witte et al. (2018) noted that SHADOZ ozone values are most 346 uncertain in the narrow region between 15 and 17 km (~3000-4300 s). However, the large 347 offsets recorded in Fig. 6b originate from four JOSIE tests conducted with TTL ozone equivalent 348 to 2 DU (e.g. Simulation 171, Fig. 5a); a value that applies to only ~ 5% of tropical SHADOZ 349 readings. Realistically, Fig. 8b in Thompson et al. (2017), based on > 6000 profiles, shows that 350 the actual TTL ozone for 12 of 14 SHADOZ stations is 8.0±1.5 DU. By 3000 s (~15 km) the 351 relative differences of all SHADOZ profiles with respect to the OPM start to increase. All 352 SHADOZ profiles show excellent agreement with OPM to within + 5% at 20-25 km (critical 353 ozone maximum). By 5000 s (~ 25 km) most SHADOZ profiles exceed OPM ozone and are 354 well-aligned with one another. The range of mean deviations in the region corresponding to 20-355 28 km is within 10%. This tighter clustering implies good measurement precision. By ~5500 s 356 (27.5 km) all the SHADOZ readings are higher than the OPM. Above 30 km the agreement 357 breaks down and there is a downturn in ozone readings relative to the OPM for most stations. 358 Exceptions are participant No. 1 and 7 that display +10% and 4% deviations, respectively (Fig. 359 **6b**). The negative relative differences are not surprising. Witte et al. (2017) showed that even

reprocessed SHADOZ ozonesonde data above ~30 km are highly variable and not as reliable.

360

361

362

363

364

365

366

367

368

369

370

371

372

373

374

375

376

377

378

379

380

381

382

383

How do column amounts for the SHADOZ participants compare on average to OPM ozone? Answers appear in **Table 5.** For the five SHADOZ simulations all of the participants record, on average, slightly more ozone than the OPM, with ratios from 1.017-1.040 (1.7% to 4.0% more O<sub>3</sub>). This result seems to validate the quality assurance practices of the SHADOZ stations, with 7 of 8 participants following the WMO-recommended instrument SST combinations and SOP (Smit et al., 2007; 2012). The segment column comparisons (columns 0-15 km, 12-18 km, 15 km-end in **Table 5**) demonstrate that the good agreement between sondes and the OPM is dominated by the ozone column from 15 km-end, i.e., the stratospheric portion of the profile. Because the WMO recommendations are largely based on JOSIE-2000, several follow-on lab tests and the BESOS conducted in 2004, it can be inferred that the WMO recommendations (Smit et al., 2012) are still valid. Agreement in the TTL (12-18 km column) averages < 0.95 for half of the groups (**Table 5**). Because the OPM recorded only 5 DU on average in this region, the larger offsets do not detract from the good agreement overall. **JOSIE SOP.** The sonde partial pressure offsets from the OPM and relative differences for the eight participants using the JOSIE 1.0.1 SST and preparation protocols appear in Fig. 7a and Fig. 7b, respectively. When these results are compared to those with the SHADOZ SOP (**Fig. 6**) two differences are observed. First, the divergence among stations is less with the more uniform specifications of the JOSIE SOP, especially in the mid-troposphere through the TTL. This is not surprising because the use of a single SST and SOP is expected to minimize variations due to SST. The JOSIE SOP uses solutions with less buffer by a factor of 2 or 10. Thus, due to the lower buffer the sonde responses show less hysteresis effect in the region with relatively fast ozone changes, resulting in increased SNR. This is particularly true in the TTL at the tropopause and just above, corresponding to the 2500 to 3500 s region in **Fig. 6b** and **Fig. 7b**.

The second difference is that ozone readings throughout the profile are lower relative to the OPM with the JOSIE SOP than the SHADOZ SOP, particularly in the troposphere (**Fig. 7a** below 4000 s) and even more so in the stratosphere, where the offsets are -1 to -2 mPa ozone. The result is a mean sonde TCO offset with the JOSIE SOP relative to the OPM of 0.97 (first two entries in column three of **Table 6**) compared to a mean 1.03 TCO offset with the SHADOZ SOP. Background cell currents and response times improved significantly during the JOSIE SOP in both sessions when a shared zero-air system was used.

SHADOZ-JOSIE Comparisons. Fig. 8a displays the average differences between the SHADOZ and JOSIE SOP profiles for Session 1. For each participant in Session 1, five simulations were made totaling 20 profiles of each SOP, both using the same SST. Up to 10 km the SHADOZ SOP resulted in relatively higher ozone readings; toward the TTL the JOSIE SOP resulted in higher ozone readings. The stratospheric differences, however, show the JOSIE SOP averages 3% lower TCO than the OPM while the SHADOZ SOP averages 3% higher TCO than the OPM (and stratospheric segment, Table 6). Note that the near-zero simulated ozone represents a small fraction of what is observed in SHADOZ records; thus, the large uncertainties seen in Fig. 8a represent the extrema of the data set.

In Session 2, to compensate for the reduced sensitivity of the 1.0%, 1/10<sup>th</sup> Buffer SST (JOSIE 1.0.1), solutions with the JOSIE SOP were prepared with twice as much KI but the same low buffer, the so-called JOSIE 2.0.1. JOSIE 1.0.1 comparisons were all made with ENSCI, whereas the JOSIE 2.0.1 referred to a combination of SPC and ENSCI. Mean profile comparisons with the different SSTs are summarized in **Fig. 8b**. The differences are not statistically significant throughout the troposphere or TTL but the JOSIE 2.0.1 profile mean is closer to the OPM in the upper stratosphere (above 5000 s). In Session 2, the ratio of sonde to OPM partial column ozone above 20 km for JOSIE 1.0.1 was 0.95, while for JOSIE 2.0.1 it was

408 0.97. Sondes filled with both SST show sondes measure less ozone than the OPM in the
409 stratosphere and are highly variable above 30 km, consistent with Fig. 7 and Witte et al. (2018)
410 findings.

Previous JOSIE campaigns and various field tests (especially the BESOS in 2004) noted that throughout the ozone profile when the same SST is used, the ENSCI instrument tends to measure more ozone than the SPC instrument. Of the 14 SHADOZ stations, 11 use the ENSCI instrument and three use the SPC type (Thompson et al., 2017; Witte et al., 2017; 2018). **Fig. 8c,** based on the combined session simulations (JOSIE 1.0.1), shows that, also for the less buffered solutions, the ENSCI instrument measures slightly higher ozone than the SPC with the greatest discrepancies in the troposphere, consistent with previous JOSIE studies.

#### **Conclusions**

- 1. All 8 stations participating in JOSIE-SHADOZ-2017 measured ozone that agreed well with the OPM.
- The slight ENSCI SPC ozone bias (ENSCI reads higher) previously observed (Smit et
   al., 2007; 2012) remained in JOSIE-SHADOZ 2017.
  - 3. JOSIE-2017 affirms the very high quality of the SHADOZ methods that use SOP and SST-instrument combinations based on earlier JOSIE campaigns and field tests as summarized in Smit et al. (2007; 2012). This is independent confirmation of the accuracy of the large SHADOZ dataset that up to now has only been compared to data from satellite and ground-based instruments (Thompson et al., 2017; Witte et al., 2017). The ozonesonde community goals of "5% accuracy and precision in TCO" has been met by SHADOZ operators engaging in collaborative ozonesonde "expert" activities since 2000. Except for the TTL, most instrument-SST combinations tested in JOSIE with

SHADOZ SOP agreed within 3% of OPM in total column amount (sonde higher) and 5-
10% throughout the ozone profile. The often large TTL ozone underestimate (>30%
relative to OPM in some tests) contributes only 2-3% of the total ozone column.

4. JOSIE tested solutions with a reduced buffer SST, of the type used at four SHADOZ stations. As expected, agreement of sonde ozone data with the OPM in the TTL regions was improved. However, sensitivity to stratospheric ozone is reduced, so TCO from these tests averaged 3% lower than the OPM. The low-bias is reduced when the KI is doubled (JOSIE2.0.1). However, the divergence of profiles with the different SST is so small (~5%) that further analysis, such as taking into account individual sonde responses, is required.

#### 5. JOSIE SOP:

- Lower, uniform, and better reproducible background cell currents are achieved using a high quality no-ozone filter source or purified air.
- The hysteresis effect ('memory' effect due to the buffering of the solution) is minimized which may improve the response of the sonde, particularly in the TTL where sharp ozone gradients are measured.

Because SHADOZ represents virtually all current ECC sonde practices used by the global ozone community, these findings and any SOP recommendations that ozonesonde "experts" consider in light of JOSIE-2017 should be universally valid for ECC instruments.

Establishing SOP guidelines and standardization of ground equipment is essential to achieving an uncertainty less than 5% between surface and 30 km altitude. The JOSIE-SHADOZ 2017 experience highlights the necessity of having a continuous reference calibration facility (WCCOS) operating over the past 25 years. The capacity building exercise has empowered

456	participants to continue working towards ensuring high quality standard in sonde data-taking.
457	With well-trained and motivated operators, SOPs based on best practices, and experiments such
458	as JOSIE-SHADOZ, our aim of an uncertainty less than 5% can be achieved.
459	
460	<b>Acknowledgments.</b> The 2017 JOSIE would not have been possible without strong support of
461	the IEK-8 group of Forschungszentrum-Jülich including a number of technicians and Ms. Gabi
462	Nork who was in charge of the organizing. The United Nations Environmental Programme (S.
463	Mylona, Nairobi) provided travel funds for SHADOZ operators. Support was also given by the
464	Science Pump Corporation (S. Schwartz), ENSCI (J. Harnetiaux and G. Kok) and JAMSTEC (S-Y.
465	Ogino). SHADOZ is supported by NASA's Upper Atmosphere Research Program and Aura
466	Validation, NOAA's Global Monitoring Division Lab (Boulder) and more than 15 organizations in
467	SHADOZ host nations and from Japan, Switzerland, France and the Netherlands.
468	
469	ACRONYMS
470	ASOPOS = Assessment of SOP for Ozone Sondes
471	DAS = Data Acquisition System
472	ECC = Electrochemical Concentration Cell
473	ENSCI = Environmental Science Corp.
474	ESC = Environmental Simulation Chamber
475	FZJ = Forschungscentrum - Jülich
476	GAW = Global Atmospheric Watch
477	GSFC = Goddard Space Flight Center
478	IOC = International Ozonesonde Commission
479	JAMSTEC = Japan Agency for Marine-Earth Science and Technology
480	JOSIE = Jülich Ozonesonde Intercomparison Experiment

- 481 KNMI = Koninklijk Nederlands Meteorologisch Instituut
   482 MLS = Microwave Limb Sounder
- NDACC = Network for the Detection of Atmospheric Composition Change
- 484 OPM = Ozone Photometer
- 485 OPS = Ozone Profile Simulator
- 486 PEF = Pump Efficiency Factor
- 487 QA = Quality Assurance
- 488 SBUV = Solar Backscatter Ultraviolet
- 489 SHADOZ = Southern Hemisphere Additional Ozonesondes
- 490 SNR = Signal Noise Ratio
- 491 SOP = Standard Operating Procedures
- 492 SPARC = Stratospheric Processes And their Role in Climate
- 493 SPC = Science Pump Corporation
- 494 SST = Sensing Solution Type
- 495 TCO = Total Column Ozone
- 496 TTL= Tropical Tropopause Layer (or Tropopause Transition Layer)
- 497 UNEP = United Nations Environmental Programme
- 498 WMO = World Meteorological Organization
- 499 WCCOS = World Calibration Centre for Ozonesondes

#### 501 **REFERENCES**

- Deshler, T., J. Mercer, H. G. J. Smit, R. Stubi, G. Levrat, B. J. Johnson, S. J. Oltmans, R. Kivi, J.
- Davies, A. M. Thompson, J. Witte, F. J. Schmidlin, G. Brothers and T. Sasaki (2007),
- Atmospheric comparison of electrochemical cell ozonesondes from different manufacturers,

- and with different cathode solution strengths: The Balloon Experiment on Standards for
- 506 Ozonesondes, J. Geophys. Res., 113, D04307, doi: 10.1029/2007JD008975.
- Deshler, T., R. Stübi, F. J. Schmidlin, J. L. Mercer, H. G. J. Smit, B. J. Johnson, R. Kivi and B.
- Nardi (2017), Methods to homogenize electrochemical concentration cell (ECC) ozonesonde
- measurements across changes in sensing solution concentration or ozonesonde manufacturer,
- 510 Atmos. Meas. Tech., 10, 2021-2043, doi:10.5194/amt-10-2021-2017.
- Fishman, J., K.W. Bowman, J. P. Burrows, A. Richter, K. V. Chance, D. P. Edwards, R. V.
- Martin, G. A. Morris, R. B. Pierce, J. R. Ziemke, J. A. Al-Saadi, J. K. Creilson, T. K.
- Schaack, and A.M. Thompson (2008), Remote sensing of tropospheric pollution from space,
- 514 Bull. Amer. Meteor. Soc., 89, 805-821.
- Hilsenrath, E., et al (1986), Results from the balloon ozone intercomparison campaign (BOIC),
- 516 J. Geophys. Res., 91 (D12), 13137–13152, doi: 10.1029/JD091iD12p13137.
- Hubert, D., et al., (2016), Ground-based assessment of the bias and long-term stability of 14 limb
- and occultation ozone profile data records, *Atmos. Meas. Tech.*, 9, 2497-2534,
- 519 doi:10.5194/amt-9-2497-2016, 2016.
- Jensen, A. A., A. M. Thompson, and F. J. Schmidlin (2012), Classification of Ascension Island
- and Natal ozonesondes using self-organizing maps, *J. Geophys. Res.*, 117, D04302, doi:
- 522 10.1029/2011JD016573.
- Johnson, B. J., S. J. Oltmans, H. Voemel, H. G. J. Smit, T. Deshler, and C. Kroeger (2002), ECC
- Ozonesonde pump efficiency measurements and tests on the sensitivity to ozone of buffered
- and unbuffered ECC sensor cathode solutions, *J. Geophys. Res.*, 107, D19 doi:
- 526 10.1029/2001JD000557.

527 Kley, D., P. J. Crutzen, H. G. J. Smit, H. Vömel, S. J. Oltmans, H. Grassl, and V. Ramanathan, 528 (1996). Observations of near-zero ozone concentrations over the convective Pacific: Effects 529 on air chemistry. Science, 274(5285), 230-233, doi: 10.1126/science.274.5285.230. 530 Komhyr, W.D. (1986), Operations handbook - Ozone measurements to 40 km altitude with 531 model 4A-ECC-ozone sondes, NOAA Techn. Memorandum ERL-ARL-149. 532 Komhyr, W. D., R. A. Barnes, G. B. Brothers, J. A. Lathrop and D. P. Opperman (1995), 533 Electrochemical concentration cell ozonesonde performance evaluation during STOIC 1989, 534 J. Geophys. Res., 100(5) 9231-9244, doi:10.1029/94JD02175. 535 McPeters, R. D., and G. J. Labow (2012) Climatology 2011: An MLS and sonde derived ozone 536 climatology for satellite retrieval algorithms, J. Geophys. Res., 117, D10303, doi: 537 10.1029/2011JD017006. 538 Melamed, M. L., P. S. Monks, A. H. Goldstein, M. G. Lawrence and J. Jennings (2016) The international global atmospheric chemistry (IGAC) project: Facilitating atmospheric 539 540 chemistry research for 25 years, Anthropocene, <a href="http://dx.doi.org/10.1016/">http://dx.doi.org/10.1016/</a> 541 j.ancene.2015.10.001 542 Mohnen, V. (1996), Update on IGAC's Global Tropospheric Ozone Network (GLONET) 543 Activity and the International Tropospheric Ozone Years (ITOY), IGACtivities, IGAC 544 Project Office, Boulder, CO, 4-6, (http://igacproject.org/sites/default/files/2016-545 07/Issue 03 Jan 1996.pdf) 546 Newton, R., G. Vaughan, H. M. A. Ricketts, L. L. Pan, A. J. Weinheimer and C. Chemel (2016), 547 Ozonesonde profiles from the West Pacific Warm Pool: measurements and validation, 548 Atmos. Chem. Phys., 16, 619-634, doi:10.5194/acp-16-619-2016.

Proffitt, M.H., and R.J. McLaughlin (1983), Fast response dual-beam UV-absorption photometer 549 550 suitable for use on stratospheric balloons, Rev. Sci. Instrum., 54, 1719-1728. 551 Rex, M., I. Wohltmann, T. Ridder, R. Lehmann, K. Rosenlof, P. Wennberg, D. Weisenstein, J. 552 Notholt, K. Krüger, V. Mohr and S. Tegtmeier (2014), A tropical West Pacific OH minimum 553 and implications for stratospheric composition, Atmos. Chem. Phys., 14, 4827-4841, 554 https://doi.org/10.5194/acp-14-4827-2014. 555 Smit H. G. J and D. Kley (1998), JOSIE: The 1996 WMO International intercomparison of 556 ozonesondes under quasi flight conditions in the environmental simulation chamber at Jülich, 557 WMO Global Atmosphere Watch Report No. 130, WMO TD No. 926, World Meteorological 558 Organization, Geneva. 559 Smit H. G. J., W. Sträter, M. Helten and D. Kley (2000), Environmental simulation facility to 560 calibrate airborne ozone and humidity sensors. Jül Berichte Nr 3796, Forschungszentrum 561 Jülich, Germany. 562 Smit, H. G. J., W. Straeter, B. Johnson, S. Oltmans, J. Davies, D. W. Tarasick, B. Hoegger, R. 563 Stubi, F. Schmidlin, T. Northam, A. M. Thompson, J. Witte, I. Boyd and F. Posny (2007), 564 Assessment of the performance of ECC-ozonesondes under quasi-flight conditions in the 565 environmental simulation chamber: Insights from the Jülich Ozone Sonde Intercomparison Experiment (JOSIE), J. Geophys. Res., 112, D19306, doi:10.1029/2006JD007308. 566 567 Smit, H. G. J., and O3S-DQA (2012), Guidelines for homogenization of ozonesonde data, 568 SI2N/O3S-DQA activity as part of "Past changes in the vertical distribution of ozone 569 assessment", available at: http://wwwdas.uwyo.edu/%7Edeshler/NDACC\_O3Sondes/O3s\_DQA/O3S-DQA-570 571 Guidelines%20Homogenization-V2-19November2012.pdf.

5/2	Smit, H. G. J., and the Panel for the Assessment of Standard Operating Procedures for
573	Ozonesondes (ASOPOS) (2014), Quality assurance and quality control for ozonesonde
574	measurements in GAW, World Meteorological Organization, GAW Report 201. [Available
575	at
576	http://www.wmo.int/pages/prog/arep/gaw/documents/FINAL_GAW_201_Oct_2014.pdf.]SP
577	ARC-IOC-GAW Assessment of Trends in the Vertical Distribution of Ozone (1998), SPARC
578	Report No.1, WMO Global Ozone Research and Monitoring Project Report No. 43, World
579	Meteorological Organization, Geneva.
580	Steinbrecht, W., et al. (2017) An update on ozone profile trends for the period 2000 to 2016,
581	Atmos. Chem. Phys., 17, 10675-10690, https://doi.org/10.5194/acp-17-10675-2017.
582	Sterling, C. W., B. J. Johnson, S. J., Oltmans, H. J. G. Smit, A. F. Jordan, P. D. Cullis, E. G.
583	Hall, A. M. Thompson, and J. C. Witte (2018), Homogenizing and estimating the uncertainty
584	in NOAA's long-term vertical ozone profile records measured with the electrochemical
585	concentration cell ozonesonde, Atmos. Meas. Tech., 11, 3661-3687,
586	https://doi.org/10.5194/amt-11-3661-2018.
587	Tarasick, D. W., et al. (2018), Tropospheric Ozone Assessment Report: Tropospheric ozone
588	observations – How well do we know tropospheric ozone changes? Elementa-Sci.
589	Anthropocene, submitted, 2018.
590	Thompson, A. M. (1992), The oxidizing capacity of the Earth's atmosphere: Probable past and
591	future changes. Science 256 1157-1165 [doi: 10.1126/science.256.5060.1157]
592	Thompson, A. M., K. E. Pickering, D. P. McNamara, M. R. Schoeberl, R. D. Hudson, J. H. Kim,
593	E. V. Browell, V. W. J. H. Kirchhoff and D. Nganga (1996) Where did tropospheric ozone

594 over southern Africa and the tropical Atlantic come from in October 1992? Insights from 595 TOMS, GTE/TRACE-A and SAFARI-92, J. Geophys. Res., 101, 24,251-24,278. 596 Thompson, A. M., et al. (2003a), Southern Hemisphere Additional Ozonesondes (SHADOZ) 597 1998–2000 tropical ozone climatology 1. Comparison with Total Ozone Mapping 598 Spectrometer (TOMS) and ground-based measurements, J. Geophys. Res., 108, 8238, 599 doi:10.1029/2001JD000967. 600 Thompson, A. M., et al. (2003b), Southern Hemisphere Additional Ozonesondes (SHADOZ) 601 1998–2000 tropical ozone climatology 2. Tropospheric variability and the zonal wave-one, J. 602 Geophys. Res., 108, 8241, doi:10.1029/2002JD002241. 603 Thompson, A. M., J. C. Witte, S. J. Oltmans and F. J. Schmidlin (2004), SHADOZ (Southern 604 Hemisphere ADditional Ozonesondes): A tropical ozonesonde-radiosonde network for the 605 atmospheric community, Bull. Amer. Meteorol. Soc., 85, 1549-1564, Thompson, A. M., J. C. Witte, H. G. J. Smit, S. J. Oltmans, B. J. Johnson, V. W. J. H. Kirchhoff, 606 607 and F. J. Schmidlin (2007), Southern Hemisphere Additional Ozonesondes (SHADOZ) 608 1998–2004 tropical ozone climatology: 3. Instrumentation, station-to-station variability, and 609 evaluation with simulated flight profiles, J. Geophys. Res., 112, D03304, 610 doi:10.1029/2005JD007042. 611 Thompson, A. M., S. J. Oltmans, D. W. Tarasick, P. von der Gathen, H. G. J. Smit, J. C. Witte, 612 (2011), Strategic ozone sounding networks: Review of design and accomplishments, Atmos. 613 Environ., 45, 2145-2163, doi:10.1016/j. atmosenv.2010.05.002. 614 Thompson, A. M., et al. (2012) Southern Hemisphere Additional Ozonesondes (SHADOZ) 615 tropical ozone climatology: Tropospheric and tropical tropopause layer (TTL) profiles with

616	comparisons to OMI based ozone products. J. Geophys. Res., 117, D23301, doi:
617	10.1029/2010JD016911.
618	Thompson, A. M., et al. (2017) First reprocessing of Southern Hemisphere Additional
619	Ozonesondes (SHADOZ) Ozone Profiles (1998-2016). 2. Comparisons with satellites and
620	ground-based instruments, J. Geophys. Res., 122, doi: 10.1002/2017 JD027406.
621	Witte, J. C., et al. (2017) First reprocessing of Southern Hemisphere ADditional OZonesondes
622	(SHADOZ) profile records (1998-2015) 1: Methodology and evaluation, J. Geophys. Res.,
623	122, doi: 10.1002/2016JD026403.
624	Witte, J. C., A. M. Thompson, H. G. J. Smit, H. Vömel, F. Posny and R. Stübi (2018) First
625	reprocessing of Southern Hemisphere Additional Ozonesondes (SHADOZ) Profile Records.
626	3. Uncertainty in ozone profile and total column, <i>J. Geophys. Res.</i> , 123, doi: 10.1002/
627	2017JD027791.
628	WMO/UNEP, Scientific Assessment of Ozone Depletion: 1990, Global Ozone Research and
629	Monitoring Project - Report No. 25, World Meteorological Organization, Geneva, 1991.
630	WMO/UNEP, Scientific Assessment of Ozone Depletion: 1994, Global Ozone Research and
631	Monitoring Project - Report No. 37, World Meteorological Organization, Geneva, 1995.
632	WMO/UNEP, Scientific Assessment of Ozone Depletion: 2010, Global Ozone Research and
633	Monitoring Project - Report No. 52, World Meteorological Organization, Geneva, 2011.
634	WMO/UNEP, Scientific Assessment of Ozone Depletion: 2014, Global Ozone Research and
635	Monitoring Project - Report No. 55, World Meteorological Organization, Geneva, 2015.
636	
637	Sidebar 1: Ozone in the Earth's Atmosphere

The ozone molecule (O<sub>3</sub>) plays several important roles in the earth's atmosphere. Its absorption of radiation warms the stratosphere, leading to the temperature inversion between troposphere and stratosphere (**Fig. SB1**). The inversion is typically referred to as the tropopause but we use the term "tropopause transition layer" to signify that the tropopause is a region (~130-70 hPa) in which a number of physical properties gradually change. Eighty-ninety percent of the ozone molecules reside in the stratosphere so harmful uv radiation is blocked from reaching the earth's surface. In the free troposphere, ozone acts as a greenhouse gas and is estimated to be responsible for ½ to 1/3 of earth's warming over the past 200 years. Tropospheric ozone is also a source of the OH free radical, the primary oxidant in the atmosphere, responsible for reacting with hundreds of species (Thompson, 1992). Ozone at the surface is considered a pollutant, harmful to human and plant health when it exceeds 3 mPa (**Fig. SB1**).

### Sidebar 2: Capacity building during JOSIE-SHADOZ

A unique feature of JOSIE-SHADOZ was that the ozonesondes were prepared by operators from organizations representing eight SHADOZ sites (see Fig. 2 showing group photos taken during both sessions in front of the WCCOS chamber). Capacity-building activities during both sessions included lectures on sonde quality-assurance, the importance of metadata reporting, troubleshooting, and training with coaches from sponsoring organizations: NASA/GSFC; NOAA/GMD; KNMI (Netherlands); KMI (Belgium); Meteoswiss, Environment – Climate Change Canada; the Finnish Meteorological Institute. Financial support for the tropical operators came from the UNEP-sponsored Vienna Convention Trust Fund, administered by WMO. Operators are essential contributors to ozonesonde quality assurance by providing detailed metadata information on each sonde launch and maintaining uniformity in their preparation and launch procedures. Bringing

together SHADOZ operators for training and knowledge sharing helps to ensure that best practices are applied to operations in a consistent manner across the SHADOZ network.

### Sidebar 3: Design of the ESC, Reference Instrument, Data System.

The WCCOS, the only one of its kind, was established in the mid-1990s at FZ-Jülich to test, calibrate and compare different types of balloon borne ozonesondes that are used to measure the distribution of ozone in the troposphere and lower/middle stratosphere. The facility is described in more detail in Smit et al. (2000): http://www.fz-juelich.de/iek/iek-8/EN/Expertise/Infrastructure/ESF/ESF node.html.

The setup of the simulation facility (Fig. SB2a), consists of four major components:

- 1. Environmental Simulation Chamber. The ESC chamber is a temperature-controlled vacuum chamber with a test room volume of about 500 liter (80 x 80 x 80 cm). Within the ESC the pressure and temperature can be dynamically regulated, with pressures between 5 and 1000 hPa and temperatures between 200 and 300 K, with a maximum rate of  $\pm 2$ K/min. Isothermically operated, the temperature variations of the air as well as the wall inside the test room can be maintained within  $\pm$  0.2 K. For more details see Smit et al. (2000).
- 2. Ozone Photometer (OPM), Ozone reference. The OPM is a fast response dual-beam UV-absorption photometer, originally developed by Proffitt and McLaughlin (1983) for use on stratospheric balloons. The instrument was flown during Balloon Ozone Intercomparison (BOIC) missions in 1983/1984 (Hilsenrath et al., 1986); it was used in the Balloon Experiment on Standards for Ozone Sondes (BESOS) field campaign in Wyoming, in 2004 (Deshler et al., 2008). The OPM is an absolute measuring device with a 1-s response time at a sampling volume flow rate of about 8 l/min. The overall accuracy of ozone measurements

- made by the OPM is better than ±2% for simulated altitudes up to 25 km (pressures down to 25 hPa) and ±3.5 % at 30-35 km altitude (12-5 hPa). The instrument resides in a separate vacuum vessel which is connected to the ESC such that the UV-photometer has the same pressure conditions as inside the test chamber.
- 3. Ozone profile simulator (OPS). A gas-flow system that controls the ozone concentrations
  sampled by the instruments in the ESC, with a gas flow rate of 12-15 l/min. The OPS can
  simulate vertical ozone profiles between the surface and 35 km. The OPS can accommodate
  up to four ozonesondes, including the OPM (Fig. SB2b). The OPS has an option to specify
  ozone step functions or zero ozone to investigate the response time and background
  characteristics of ozonesondes.

4. <u>Data Acquisition System (DAS).</u> The entire simulation process is automated by computer control in order to have reproducible conditions with respect to the simulated pressure, temperature and ozone versus time, and for recording and storing the large variety of parameters measured during the simulation process. A special electronic interface (JOSIE/ ECC-interface) couples the ECC sonde to the DAS, transmitting cathode cell current, pump temperature, pump motor current and pump motor voltage (12V). A small variable electrical heater (0-10W) adjusts pump temperatures to values similar to actual flight temperatures.

### 1 Table List:

### 2 Table 1: JOSIE activities on ozonesonde procedures and related reports.

Campaign	Objective
JOSIE-1996 GAW Report #130	<ul> <li>Operating Procedures</li> <li>Profiling Capabilities</li> <li>Intercomparison sonde types (ECC, Brewer Mast, Meisei)</li> </ul>
JOSIE-1998 GAW Report #57	Manufacturing ECC sondes (SPC, ENSCI)
JOSIE-2000 GAW Report #158 (Smit et al., 2007)	<ul> <li>Operating Procedures</li> <li>Focus on ECC sonde         <ul> <li>Different sensing solution types</li> <li>Different manufacturers (SPC, ENSCI)</li> </ul> </li> </ul>
BESOS-2004 (Deshler et al., 2008)	<ul> <li>Operating Procedures under flight conditions</li> <li>Focus on ECC sonde         <ul> <li>Different sensing solution types</li> <li>Different manufacturers (SPC, ENSCI)</li> </ul> </li> </ul>
ASOPOS 2002-2012 GAW Report #201	Define and establish Standard Operating Procedures for ECC sondes
JOSIE-2009	Manufacturers (SPC, ENSCI)
JOSIE-2010	Refurbished sondes
O3S-DQA Guidelines Report-2012	Homogenization and Uncertainties
JOSIE-SHADOZ-2017	<ul> <li>Operating procedures</li> <li>Tropical simulations</li> <li>Different sensing solution types</li> <li>Different manufacturers (SPC, ENSCI)</li> </ul>

## 4 Table 2: SHADOZ stations operating at least 10 years between 1998 and 2017

Station	Latitude, Longitude	Current ECC Sensor	<b>Current Radiosonde</b>
Pago Pago, Am. Samoa	14.23S, 170.56W	ENSCI	iMet-1
Hilo, Hawaii	19.40N, 155.00W	ENSCI	iMet-1
San Cristobal, Galapagos,	0.92S, 89.60W	ENSCI	Vaisala RS92
Ecuador			
San Pedro, Costa Rica	9.94N, 84.04W	ENSCI	iMet-1
Paramaribo, Surinam	5.81N, 55.21W	SPC	Vaisala RS92
Ascension Is., U.K	7.98S, 14.42W	ENSCI	iMet-1
Natal, Brazil	5.42S, 35.38W	SPC	Lockheed-Martin-
			Sippican LMS6
Irene, S. Africa	25.90S, 28.22E	SPC	Vaisala RS92
Nairobi, Kenya	1.27S, 36.80E	ENSCI	Vaisala RS92
La Réunion, France	21.10S, 55.48E	ENSCI	Modem M10
Kuala Lumpur, Malaysia	2.73N, 101.70E	ENSCI	GRAW DFM-09
Hanoi, Vietnam	21.02N, 105.80E	ENSCI	Vaisala RS92
Watukosek-Java, Indonesia	7.57S, 112.65E	ENSCI	*
Suva, Fiji	18.10S, 178.40E	ENSCI	iMet-1
O . IM . DOH ROZOD	1' 1 1	1000 1000 X/ : 1	DC00 1000 2012

<sup>5 \*</sup>Operated Meisei RS II-KC79D radiosonde-ozonesonde system 1992-1999; Vaisala RS80 1998-2013.

- 7 Table 3: SHADOZ station operators and instruments tested in JOSIE. Stations 1-4
- 8 participated in Session 1 (9-20 October 2017); stations 5-8 participated in Session 2 (23
- 9 **October 3 November 2017).**

Participant	SST	Operator	Affiliation	Station					
Number									
	Session 1								
1	1.0% Full Buffer	Tshidi Machinini	South African Weather Service	Irene, South Africa					
2	1.0% Full Buffer	Francisco R. da Silva	Brazilian Space Agency	Natal, Brazil					
3	0.5% Half Buffer	Kennedy Thiong'o	Kenyan Meteorological	Nairobi, Kenya					
			Department						
4	4 0.5% Half Buffer Ernesto Corrales		University of Costa Rica	San Pedro, Costa Rica					
		Sessi	on 2						
5	5 1.0% Full Buffer George Paiman		Meteorological Service of	Paramaribo, Surinam					
			Suriname						
6	0.5% Half Buffer	Zamuna Zainal	Malaysian Meteorological	Kuala Lumpur,					
			Department	Malaysia					
7	7 0.5% Half Buffer Françoise Posny		Université La Réunion, Météo- La Réunion Is., F						
			France, CNRS						
8	0.5% Half Buffer Nguyen Thi Hoang Anh		h Vietnam Meteorological and Hanoi, Vietnam						
			Hydrological Administration						

- Table 4: Characteristics of JOSIE-SHADOZ-2017 simulations in the WCCOS chamber
- 12 with Simulation Numbers listed for the two Sessions. LT=lower troposphere, MT=mid-
- troposphere, UT=upper troposphere and LS=lower-stratosphere. All profiles simulated
- with nominal 5 m/s ascent velocity. The tropopause was located at Z=18-20 km with
- minimum temperature ~-(70-80)°C. The stratospheric profile was specified to be the same
- 16 for all simulations.

	Session 1								
Simulation Number	Troposphere Profile Type	Profile Type Index**	Specifications	ECC Procedure					
171	Recent deep convection	1	Extremely low O <sub>3</sub> values nearly uniformly up to tropopause with very steep gradient into LS	Station-supplied SST & procedures					
172	Maritime background	2	Low O <sub>3</sub> in LT, moderate O <sub>3</sub> in MT, extremely low O <sub>3</sub> in UT	Station-supplied SST & procedures					

173, 174, 175, 176*	Biomass burning	3	Enhanced O <sub>3</sub> in LT, high O <sub>3</sub> in MT, low O <sub>3</sub> in UT	Station-supplied SST & procedures
177, 178, 179, 181	Biomass burning	3	Enhanced O <sub>3</sub> in LT, high O <sub>3</sub> in MT, low O <sub>3</sub> in UT	JOSIE-supplied SST & WMO procedures
180	Maritime 2 background		Low O <sub>3</sub> in LT, moderate O <sub>3</sub> in MT, extremely low O <sub>3</sub> in UT	JOSIE-supplied SST & WMO procedures
			Session 2	
Simulation Number	Troposphere Profile Type	Profile Type Index**	Specifications	ECC Procedure
182, 183, 184,186	Biomass burning	3	Enhanced O <sub>3</sub> in LT, high O <sub>3</sub> in MT, low O <sub>3</sub> in UT	Station-supplied SST & procedures
185	Maritime background	2	Low O <sub>3</sub> in LT, moderate O <sub>3</sub> in MT, extremely low O <sub>3</sub> in UT	Station-supplied SST & procedures
187, 188, 190, 191	Biomass burning	3	Enhanced $O_3$ in LT, high $O_3$ in MT, low $O_3$ in UT	JOSIE-supplied SST & WMO procedures
189	Maritime background	2	Low ozone in LT, enhanced ozone in MT and extreme low ozone in UT	JOSIE-supplied SST & WMO procedures

<sup>\*</sup> Due to a problem with the ESC, Simulation 176 only recorded profiles to 15 km.

## 20 Table 5: Total and partial column statistics from two SHADOZ simulations and means for

## 21 all 10 simulations (five each in Sessions 1 and 2). All simulations use SHADOZ SOPs.

Instrument	Sim 171	Sim 182	Mean	Mean	Mean	Mean
	(DU)	(DU)	OPM/Sonde	OPM/Sonde	OPM/Sonde	OPM/Sonde
			Ratio: TCO	Ratio: Trop	Ratio: TTL	Ratio: Strat O <sub>3</sub>
				$O_3$ (0-15 km)	O <sub>3</sub> (12-18	(15 km-end)
					km)	
OPM	282		337 DU	47.0 DU	4.93 DU	298 DU
Participant 1	1.07		1.03	1.09	1.02	1.04
Participant 2	1.09		1.04	1.09	1.03	1.04
Participant 3	1.03		1.03	1.02	0.95	1.03
Participant 4	1.01		1.02	1.06	1.01	1.02
OPM		334	313 DU	41.0 DU	5.30 DU	271 DU
Participant 5		1.00	1.03	0.85	0.77	1.03
Participant 6		1.04	1.04	0.89	0.87	1.05
Participant 7		1.07	1.04	0.93	0.93	1.05
Participant 8		1.00	1.02	0.88	0.87	1.02

<sup>\*\*</sup> In Figure 4, 1 = blue, 2= green, 3= red

## 23 Table 6: Total and partial column statistics from profile simulations, relative to OPM,

# 24 categorized by SOP and sonde/solution types.

Methodology	No.	Mean	Mean	Mean	Mean
		Sonde/OPM	Sonde/OPM Trop	Sonde/OPM TTL	Sonde/OPM Strat
		TCO	O <sub>3</sub> (0-15 km)	O <sub>3</sub> (12-18 km)	O <sub>3</sub> (20 km-end)
SHADOZ SOP	40	1.03	1.01	0.94	1.04
JOSIE SOP	40	0.97	0.99	0.94	0.97
ENSCI 1.0%, 0.1B*	25	0.98	1.00	0.97	0.98
SPC 1.0%, 0.1B	10	0.97	0.96	0.90	0.98
ENSCI 0.5%, 0.5B	20	1.03	1.00	0.91	1.04
SPC 1.0%, 1.0B	15	1.03	1.01	0.95	1.04
ENSCI 2.0%, 0.1B	5	0.97	1.01	0.97	0.97
SPC 2.0%, 0.1B	5	0.97	0.94	0.90	0.96

\* B=Buffer

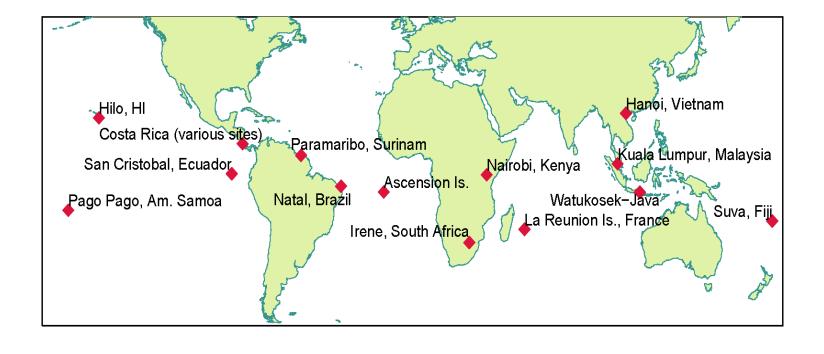
26 Figure Caption List: 27 Figure 1: Map of SHADOZ stations. 28 29 30 Figure 2(a): Session 1 participants: (1) George Brothers (NASA/WFF); (2) Kennedy 31 Thiong'o (Kenya Met Dept.); (3) Francisco Raimundo da Silva (INPE Natal); (4) Ernesto 32 Corrales (Univ. Costa Rica); (5) Peter von der Gathen (Alfred Wegener Institute); (6) Herman Smit (FZ Jülich); (7) Ryan Stauffer (NASA/GSFC); (8) Gary Morris (St. 33 34 Edward's Univ.); (9) Gabi Nork (FZ Jülich); (10) Anne Thompson (NASA/GSFC); (11) Bryan Johnson (NOAA ESRL); (12) Tshidi Machinini (South African Weather Service); 35 36 (13) Tatsumi Nakano (Japan Met Agency); (14) Rhonie Wolff (NASA/WFF). 37 Figure 2(b): Session 2 participants: (1) Gonzague Romanens (MeteoSwiss); (2) Torben 38 39 Blomel (FZ Jülich); (3) Jennifer Gläser (FZ Jülich); (4) Nguyen Thi Hoang Anh (Vietnam 40 Meteorological and Hydrological Administration); (5) Anne Thompson (NASA/GSFC); (6) Jonathan Davies (Env. Climate Change Canada); (7) Zamuna Zainal (Met Malaysia); (8) 41 42 Patrick Neis (FZ Jülich); (9) Gabi Nork (FZ Jülich); (10) Rigel Kivi (FMI); (11) Rene Stübi (MeteoSwiss); (12) Patrick Cullis (NOAA ESRL); (13) Herman Smit (FZ Jülich); (14) Marc 43 Allaart (KNMI); (15) Roeland Van Malderen (Royal Meteorological Institute of Belgium); 44 45 (16) Jacquelyn Witte (NASA/GSFC); (17) George Paiman (Met Dept. of Suriname); (18) Andreas Petzold (FZ Jülich); (19) Gilbert Levrat (MeteoSwiss); (20) Françoise Posny 46 47 (Univ. of La Réunion).

49 Figure 3: (a) Schematic of an electrochemical concentration cell (ECC) in operational 50 mode. (b) ECC instrument in Styrofoam box in which it is housed during JOSIE tests or in 51 deployment (when launched the sensor is sealed with a Styrofoam lid). Instrument and 52 solution type for each JOSIE-SHADOZ station appear in Tables 2 and 3, respectively. 53 54 Figure 4: Simulated ozone profiles (in partial pressure) as a function of simulation time for 55 the troposphere and stratosphere until 33 km altitude (a) and up to 20 km in (b). Three different tropospheric ozone profiles with extreme low ozone concentrations up to the 56 57 tropopause (Altitude ≈ 18 km) in blue and two profiles with moderate to enhanced middle 58 tropospheric ozone values in green and red, respectively. 59 60 Figure 5: Ozone "raw" profiles of typical simulations in Sessions 1 (a) and 2 (b). Participants are listed in Table 3, simulation specifications are listed in Table 4. 61 62 63 Figure 6: (a) Participant mean profiles relative to OPM in partial pressure (mPa), and (b) % deviation (Sonde – OPM / OPM). Based on 5 simulations per participant. 64 65 Figure 7: Same as Fig. 6, except for JOSIE SOP as described in Table 4. 66 67 68 Figure 8: (a) Session 1 SHADOZ SOP (blue) and JOSIE SOP (red) mean profiles subtracted from the OPM profile mean; (b) Session 2 JOSIE 2.0.1 (black) and JOSIE 1.0.1 69

(red) SST profile means subtracted from the OPM; and (c) Session 1 and 2 mean profiles of

ENSCI-OPM (red) and SPC-OPM (blue) for which JOSIE 1.0.1 SST and SOP was used. 1-71 72 sigma standard deviations for all panels are included. 73 74 Figure SB1. Ozone and temperature profiles from a typical SHADOZ sounding at Natal, 75 Brazil, taken from the archive, <a href="https://tropo.gsfc.nasa.gov/shadoz">https://tropo.gsfc.nasa.gov/shadoz</a>. 76 77 Figure SB2: (a) Set up for the simulation of vertical ozone soundings with a schematic of the Environmental Simulation Chamber, showing Ozone Photometer (OPM) standard 78 79 reference, control systems, placement of four ozonesondes ("TEO") in the chamber and data acquisition system (DAS). (b) Photo of the chamber and DAS computer. 80

81



**Fig. 1.** Map of SHADOZ stations.



Fig. 2. (a) Session 1 participants: (1) George Brothers (NASA/WFF); (2) Kennedy Thiong'o (Kenya Met Dept.); (3) Francisco Raimundo da Silva (INPE Natal); (4) Ernesto Corrales (Univ. Costa Rica); (5) Peter von der Gathen (Alfred Wegener Institute); (6) Herman Smit (FZ Jülich); (7) Ryan Stauffer (NASA/GSFC); (8) Gary Morris (St. Edward's Univ.); (9) Gabi Nork (FZ Jülich); (10) Anne Thompson (NASA/GSFC); (11) Bryan Johnson (NOAA ESRL); (12) Tshidi Machinini (South African Weather Service); (13) Tatsumi Nakano (Japan Met Agency); (14) Rhonie Wolff (NASA/WFF).

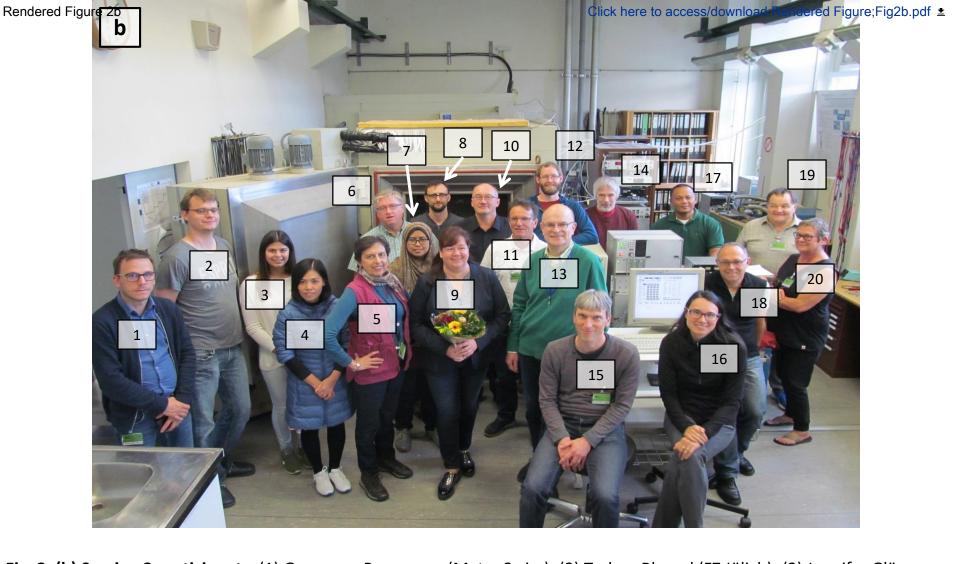
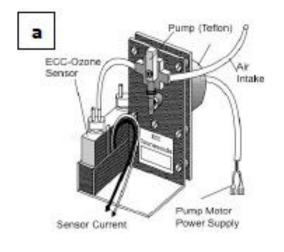


Fig. 2. (b) Session 2 participants: (1) Gonzague Romanens (MeteoSwiss); (2) Torben Blomel (FZ Jülich); (3) Jennifer Gläser (FZ Jülich); (4) Nguyen Thi Hoang Ahn (National Hydro-Meteorological Service of Vietnam); (5) Anne Thompson (NASA/GSFC); (6) Jonathan Davies (Env. Climate Change Canada); (7) Zamuna Zainal (Met Malaysia); (8) Patrick Neis (FZ Jülich); (9) Gabi Nork (FZ Jülich); (10) Rigel Kivi (FMI); (11) Rene Stübi (MeteoSwiss); (12) Patrick Cullis (NOAA ESRL); (13) Herman Smit (FZ Jülich); (14) Marc Allaart (KNMI); (15) Roeland Van Malderen (Royal Meteorological Institute of Belgium); (16) Jacquelyn Witte (NASA/GSFC); (17) George Paiman (Met Dept. of Suriname); (18) Andreas Petzold (FZ Jülich); (19) Gilbert Levrat (MeteoSwiss); (20) Françoise Posny (Univ. of La Réunion).



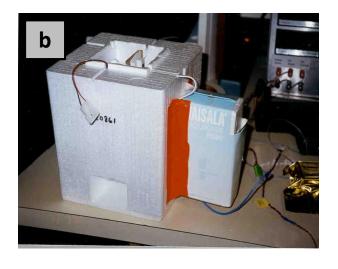
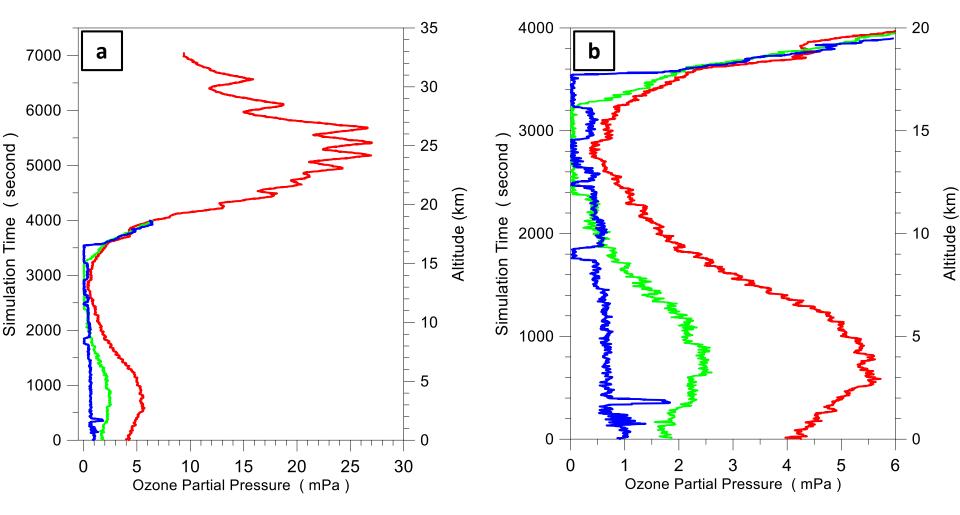
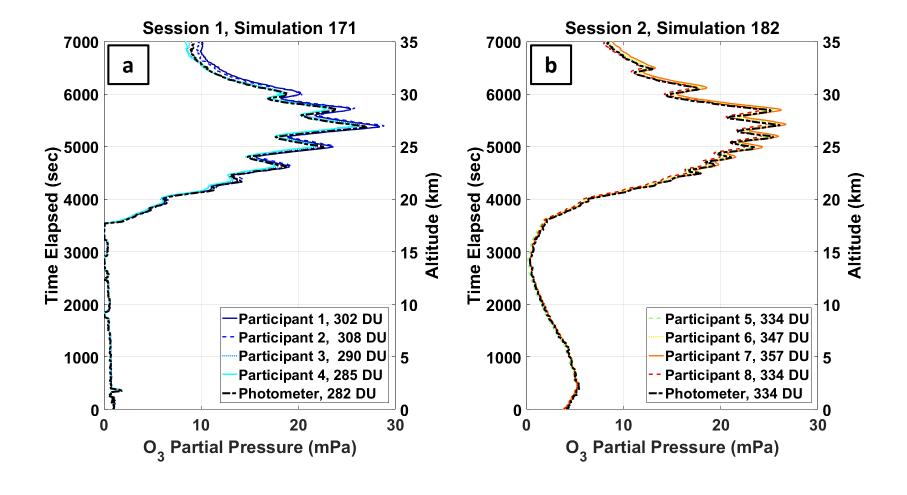


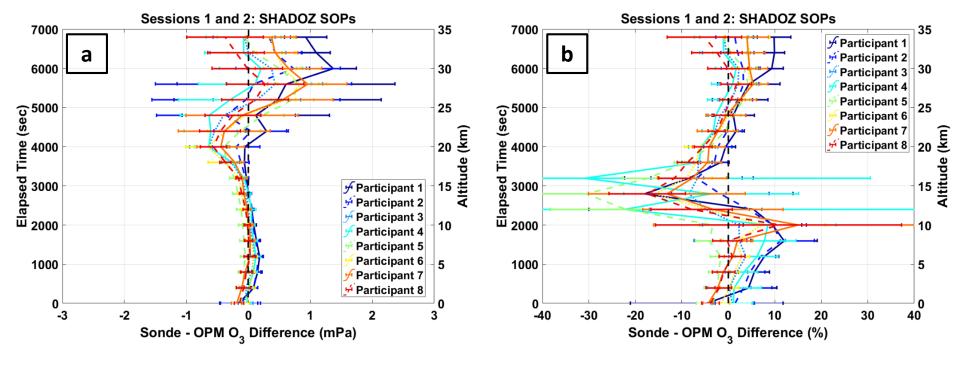
Figure 3: (a) Schematic of an electrochemical concentration cell (ECC) in operational mode. (b) ECC instrument in Styrofoam box in which it is housed during JOSIE tests or in deployment (when launched the sensor is sealed with a Styrofoam lid). Instrument and solution type for each JOSIE-SHADOZ station appear in Tables 2 and 3, respectively.



**Fig. 4.** Simulated ozone profiles (in partial pressure) as a function of simulation time for the troposphere and stratosphere until 33 km altitude (a) and up to 20 km in (b). Three different tropospheric ozone profiles with extreme low ozone concentrations up to the tropopause (Altitude  $\approx$  18 km) in blue and two profiles with moderate to enhanced middle tropospheric ozone values in green and red, respectively.



**Fig. 5.** Ozone "raw" profiles of typical simulations in Sessions 1 (a) and 2 (b). Participants are listed in Table 3, simulation specifications are listed in Table 4.



**Fig. 6.** (a) Participant mean profiles relative to OPM in partial pressure (mPa), and (b) % deviation (Sonde – OPM / OPM). Based on 5 simulations per participant.

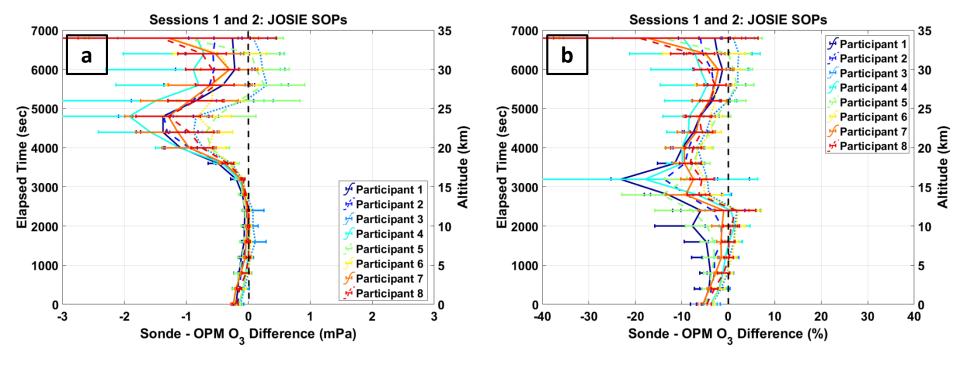
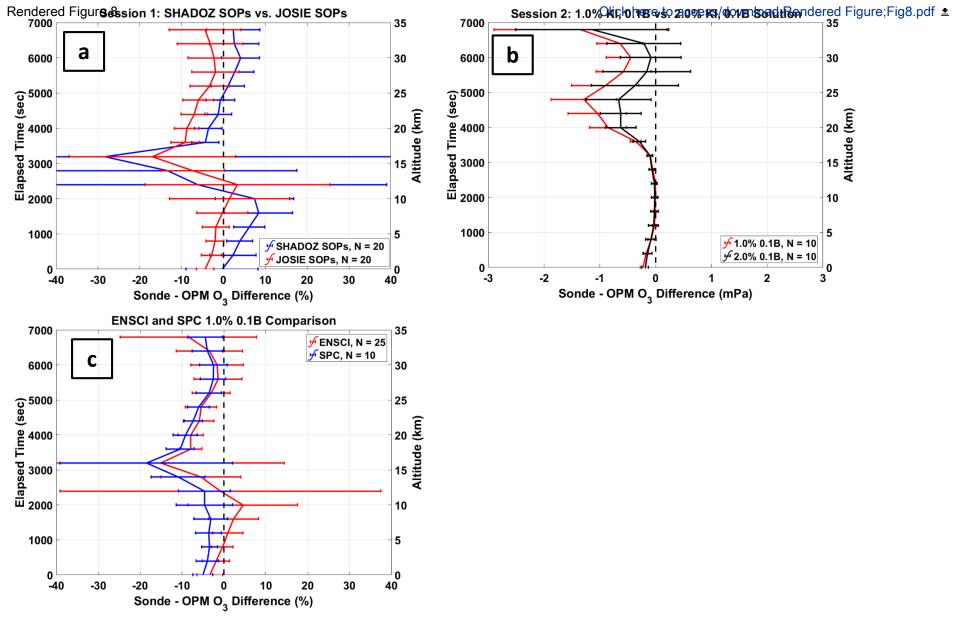


Fig. 7. Same as Fig. 6, except for JOSIE SOP as described in Table 4.



**Fig. 8.** (a) Session 1 SHADOZ SOP (blue) and JOSIE SOP (red) mean profiles subtracted from the OPM profile mean; (b) Session 2 JOSIE 2.0.1 (black) and JOSIE 1.0.1 (red) SST profile means subtracted from the OPM; and (c) Session 1 and 2 mean profiles of ENSCI-OPM (red) and SPC-OPM (blue) for which JOSIE 1.0.1 SST and SOP was used. 1-sigma standard deviations for all panels are included.

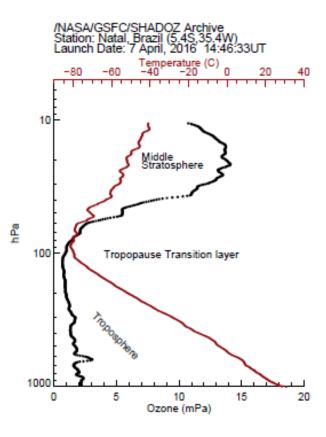
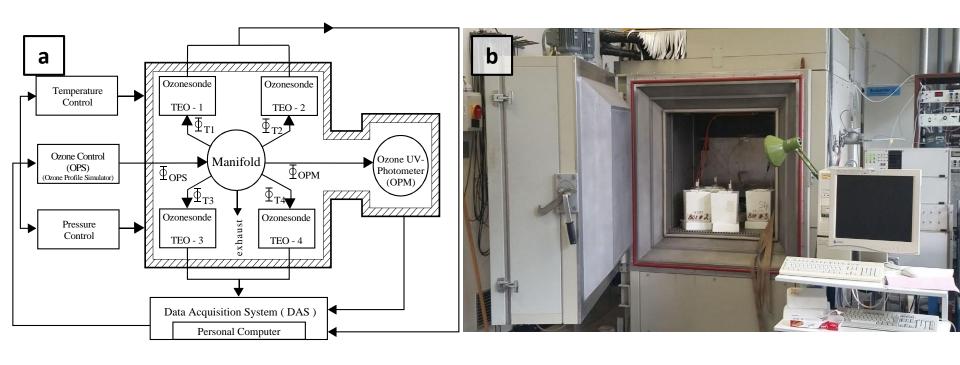


Fig. SB1. Ozone and temperature profiles from a typical SHADOZ sounding at Natal, Brazil, taken from the archive, <a href="https://tropo.gsfc.nasa.gov/shadoz">https://tropo.gsfc.nasa.gov/shadoz</a>.



**SB2.** (a) Set up for the simulation of vertical ozone soundings with a schematic of the Environmental Simulation Chamber, showing Ozone Photometer (OPM) standard reference, control systems, placement of four ozonesondes in the chamber ("TEO") and data acquisition system (DAS). (b) Photo of the chamber and DAS computer.